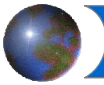


Paleoclimatology

Anna Klene
Department of Geography
University of Montana



3 Objectives

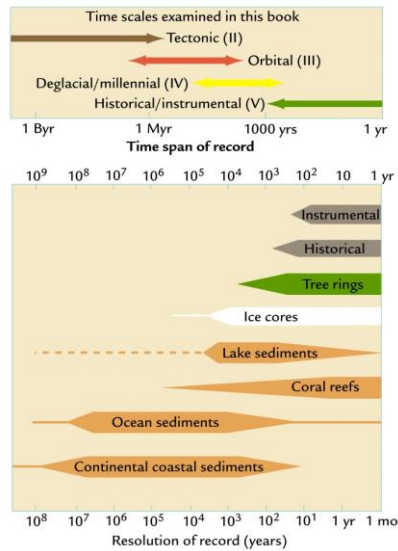
- Discuss climate archives
 - ☒ Piecing the puzzle together

- Discuss key climate events using these different archives
 - ☒ Current understanding of atm evolution

- Review key time periods of interest to current warming



Time scales for Proxy Data



Ruddiman, 2008



Archives of Climate Change:

Geological

Biological: Fossils & Pollen

Cryological: Ice Cores

Historical

Biological: Tree-Rings

Instrumental Records

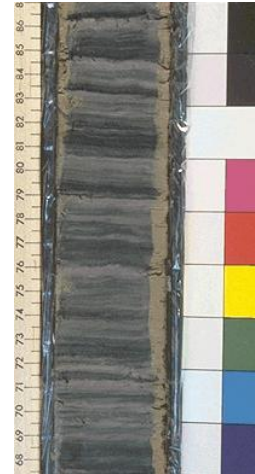
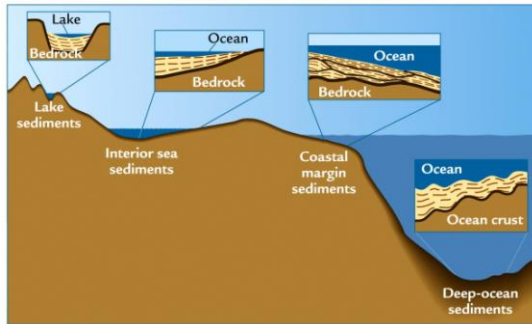
- Proxy: Using one thing in place of another...
- Always better if 2 different, independent proxies agree ☺



Archives of Climate Change:

Geological

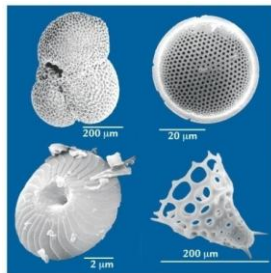
- Sediment structures & material (loess)
- Glacial moraines
- Lake sediments
- Coastal & Deep Ocean sediments



Archives:

Biological

- Fossils or dead material
- Trees
- Critters (macro: mammals, beetles, etc. & micro: corals, plankton, foraminifera, etc.)

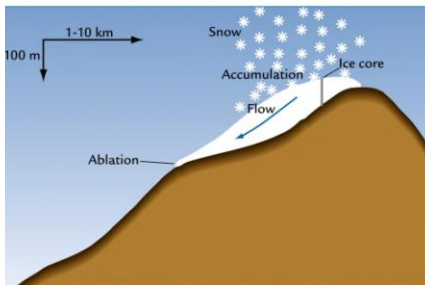




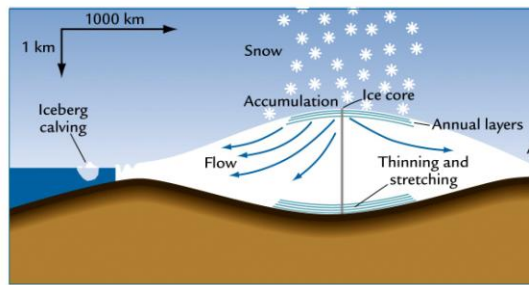
Archives of Climate Change:

Cryological

Glaciers & Ice Caps



Mountain glaciers



B

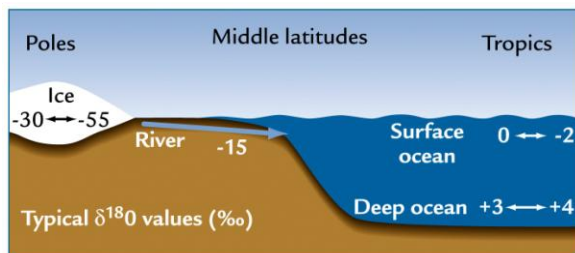
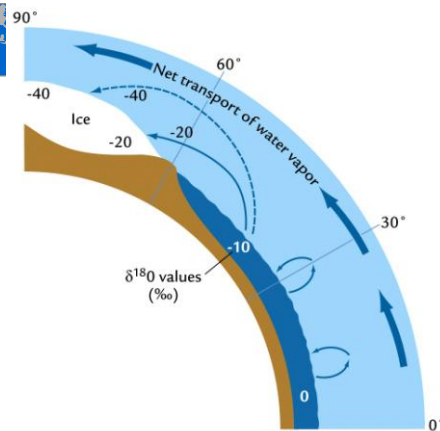
Continental ice sheets



Ice & Sediment Cores

● Oxygen-isotope analysis:

- ❑ $\delta^{18}\text{O}$
- ❑ Measure ratio of ^{16}O to ^{18}O
- ❑ Water from ocean enriched in 18 as 16 evaporates better...
- ❑ When glaciers advance, more 16 frozen, so even more 18 in water...





Ice Cores & Sediment Cores

- deuterium/hydrogen ratio:
 - ✘ $\delta D\text{‰}$
 - ✘ Measure ratio of ^2H to ^1H ...
 - ✘ Deuterium is heavier than normal Hydrogen, so it takes more energy to evaporate any water molecule made with “heavy hydrogen”.
 - ✘ The result is that the colder it gets, the less Deuterium ends up in precipitation.
 - ✘ The smaller the D/H ratio, the colder the climate.



Ice Cores

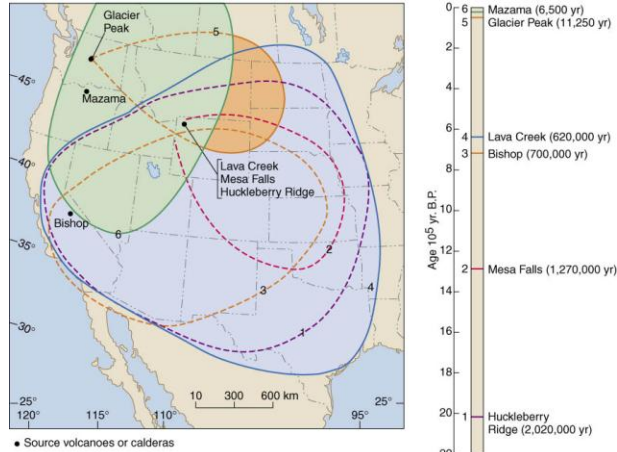
- Ice cores:
 - ✘ volcanic ash
 - ✘ particulates (dust),
 - ✘ pollen,
 - ✘ chemical composition of the air trapped inside,
 - ✘ etc..





Archives

- Volcanic Ash
- Source by chemical signature
- Provides a calibration layer across variety of deposits



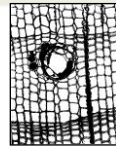
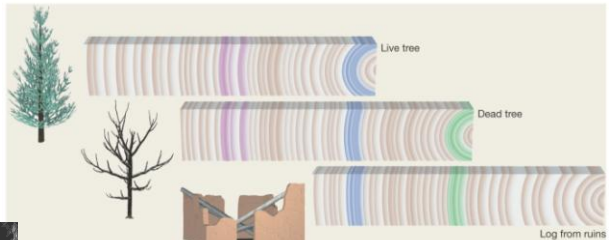
From Skinner



Tree-rings:



- Annual layers of growth
 - ☒ Depends on temp, precip, evapotrans.
 - ☒ Varies from species to species



<http://web.utk.edu/~grissino/gallery.htm>



Archives of Climate Change:

Historical Records

Letters, Diaries, Other Records

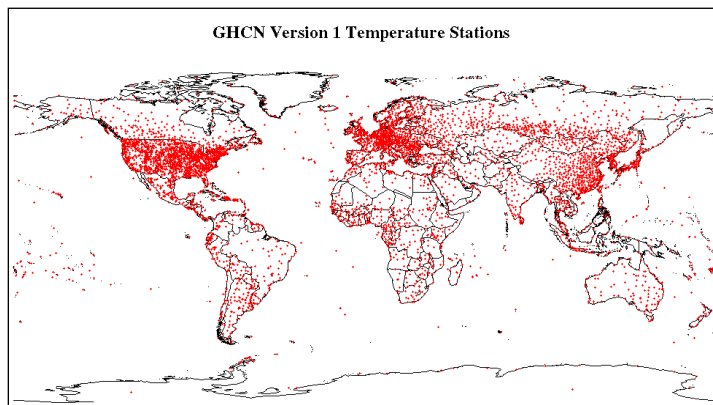
- **Hunters in the Snow**, 1565
Pieter Bruegel the Elder
(Netherlandish, ca. 1525/30—1569)
Oil on panel; 46 1/8 x 63 7/8 in. (117 x 162 cm)
Image courtesy of the Kunsthistorisches
Museum, Vienna

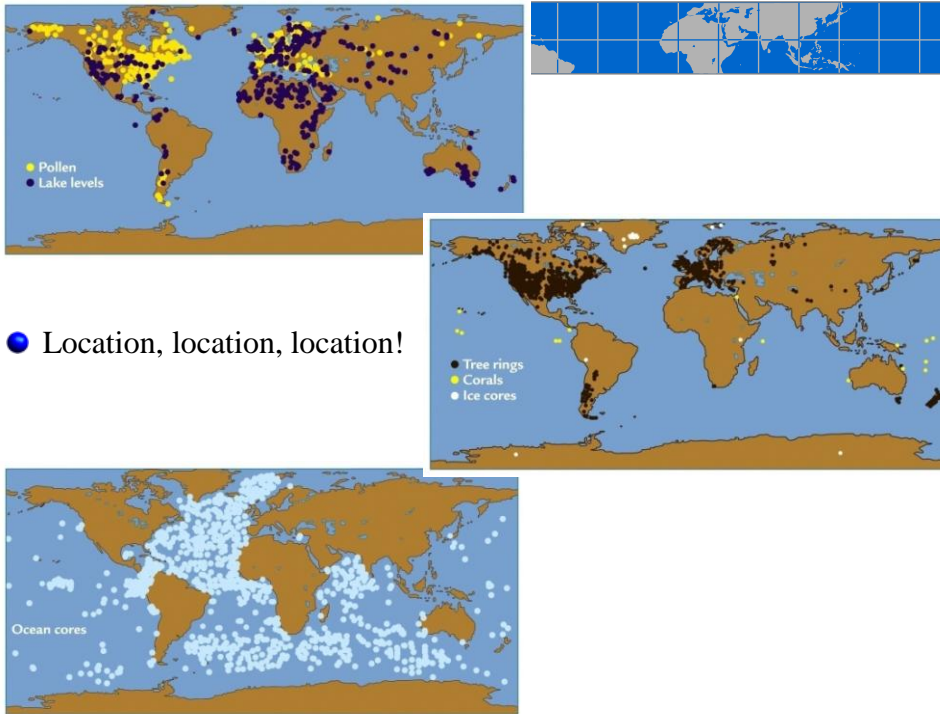


Archives of Climate Change:

Instrumental Records

Only within last ~200 years





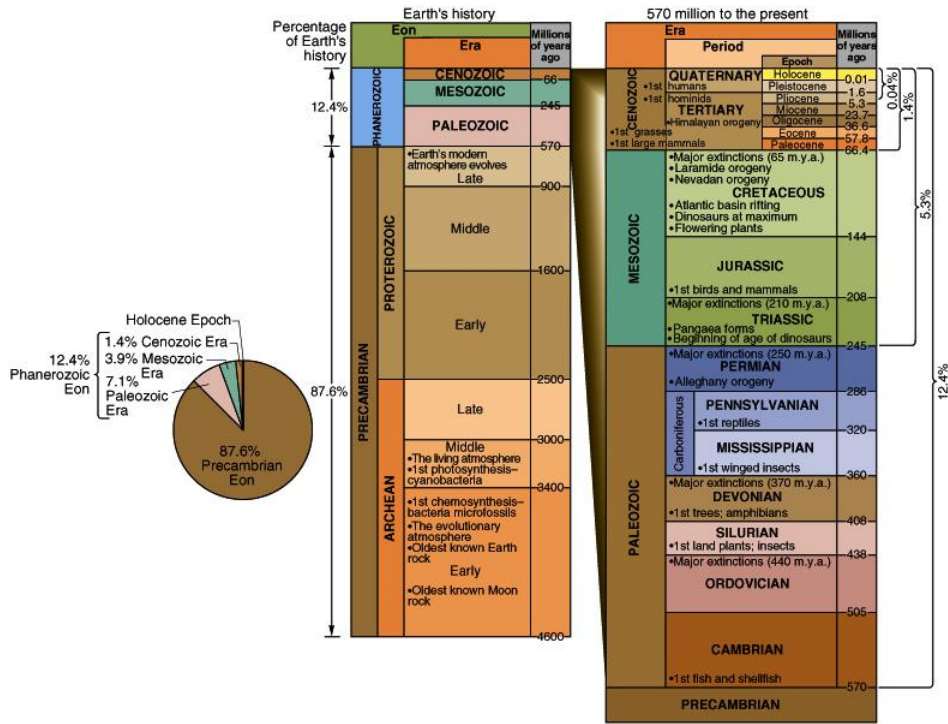
● Location, location, location!



Earth's Evolution

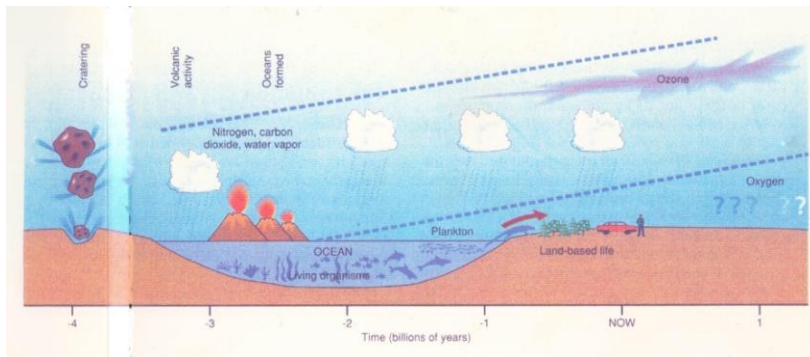
- ~4+ BYA: All blown away
- ~4: Magnetic field forms & atm held in place – no O₂
- ~3.8: Out-gassing continues but liquid earth possible as planet cools below 100 C
- ~3.5 BYA: First life forms release O₂
- ~500 MYA: O₂ levels high enough for ozone layer & plants & animals can now colonize land

● **All from geological evidence!!**



Earth's Primordial Atmosphere

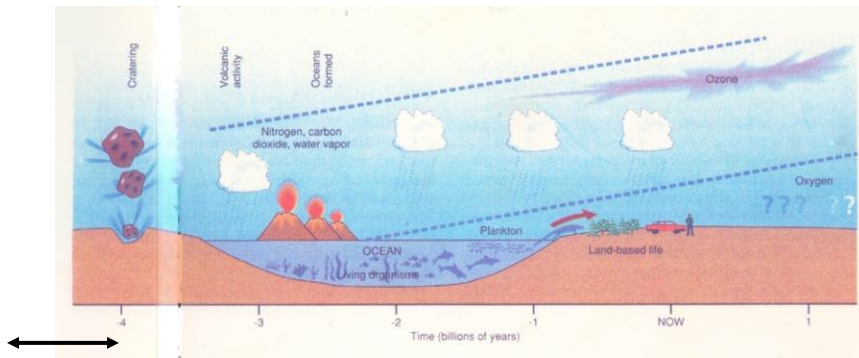
- 4+ billion years ago (Y.A.)
- Consisted of gases most abundant in solar system, hydrogen and helium (lightest elements)
- Mainly blown away





Earth's Primordial Atmosphere

- 4+ Billion Y.A.
- Begins build up once magnetic field developed
- Consists of CO₂, NO*, H₂O



Stromatolites

Photosynthesis

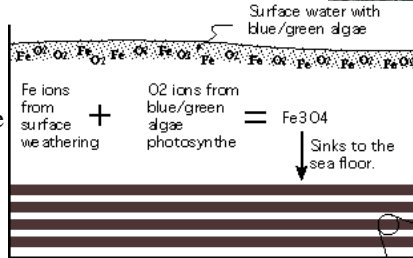
- $\text{CO}_2 + \text{H}_2\text{O} + \text{light} \rightarrow \text{CH}_2\text{O} + \text{O}_2$
- Cyanobacteria (Eubacteria) aka blue-green algae, appear ~ 3.5 bya
 - ◆ Release O₂ as byproduct
- Accumulation of O₂ in the atmosphere didn't start until oceanic Fe₂₊ was oxidized (~2 bya).



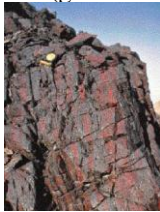


Banded Iron Formations

- Water with O₂ (from blue-green algae) & Fe from surface weathering.
- Get deposits (iron-rich layer) This cleans algae's environment.
- Too much algae, produce too much O₂, not enough Fe to remove it...
- O₂ toxic to algae, population collapse... (get white layer)



After combining the Fe and O₂ ions into Magnetite (Fe₃O₄), the mineral grains sink to the sea floor, where they accumulate into iron-rich and iron-poor layers.



The red bands are hematite, and are interbedded with chert.

In an ideal setting, you would expect the magnetite-rich layers to exhibit a reversed graded bedding. Looking from the bottom up, this would involve a slow transition into the magnetite-rich layers, representing slowly increasing O₂ levels in the upper sea water in response to the increasing population of blue/green algae. The upper contact of each magnetite-rich layer would be relatively abrupt, reflecting the sudden extinction of the population due to O₂ poisoning, and the resulting loss of available O₂ in the water to combine with the iron ions.

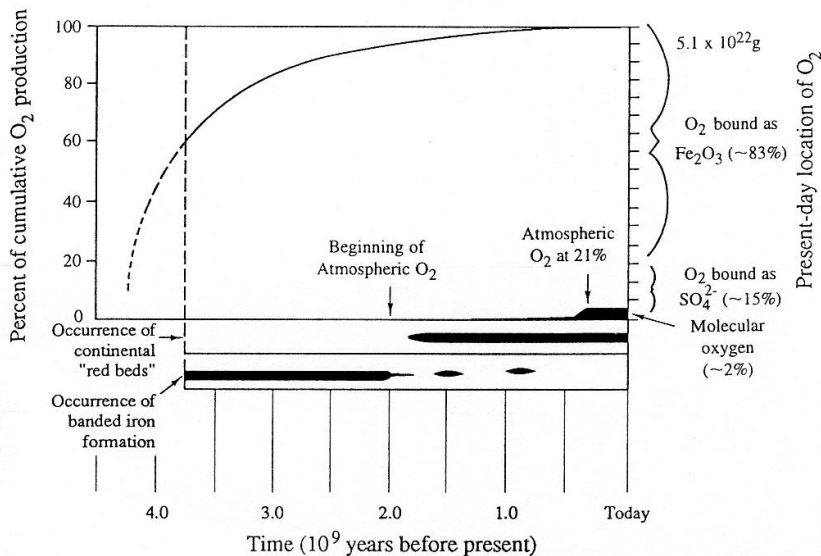
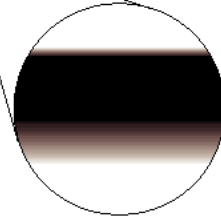


Figure 2.7 Cumulative history of O₂ released by photosynthesis through geologic time. Of more than 5.1×10^{22} g of O₂ released, about 98% is contained in seawater and sedimentary rocks, beginning with the occurrence of Banded Iron Formations at least 3.5 billion years ago (bya). Although O₂ was released to the atmosphere beginning about 2.0 bya, it was consumed in terrestrial weathering processes to form Red Beds, so that the accumulation of O₂ to present levels in the atmosphere was delayed to 400 mya. Modified from Schidlowski (1980).